

Detection of Ionized Calcium in the Atmosphere of the Ultra-hot Jupiter KELT-9b

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Abstract

With a dayside temperature in excess of 4500 K, comparable to a mid-K-type star, KELT-9b is the hottest planet known. Its extreme temperature makes KELT-9b a particularly interesting test bed for investigating the nature and diversity of gas giant planets. We observed the transit of KELT-9b at high spectral resolution ($R \sim 94,600$) with the CARMENES instrument on the Calar Alto 3.5 m telescope. Using these data, we detect for the first time ionized calcium (Ca II triplet) absorption in the atmosphere of KELT-9b; this is the second time that Ca II has been observed in a hot Jupiter. Our observations also reveal prominent H α absorption, confirming the presence of an extended hydrogen envelope around KELT-9b. We compare our detections with an atmospheric model and find that all four lines form between atmospheric temperatures of 6100 and 8000 K and that the Ca II lines form at pressures between 50 and 100 nbar while the H α line forms at a lower pressure (~10 nbar), higher up in the atmosphere. The altitude that the core of H α line forms is found to be ~1.4 R_p , well within the planetary Roche lobe (~1.9 R_p). Therefore, rather than probing the escaping upper atmosphere directly, the H α line and the other observed Balmer and metal lines serve as atmospheric thermometers enabling us to probe the planet's temperature profile, thus the energy budget.

Unified Astronomy Thesaurus concepts: Exoplanet atmospheres (487); Planetary atmospheres (1244); Extrasolar gas giants (509); Exoplanets (498); Hot Jupiters (753); Exoplanet atmospheric composition (2021)

1. Introduction

Characterizing exoplanetary atmospheres provides valuable insights into their composition, climate, evolution, and habitability. Transiting exoplanets are of particular interest because they offer the opportunity to obtain atmospheric transmission spectra (Charbonneau et al. 2000). Transit observations have been conducted in both broadband spectral windows (e.g., Sing et al. 2016) and at high spectral resolution $(R \sim 50,000-100,000; e.g., Snellen et al. 2010)$. Observing such a spectrum makes it possible to detect individual atomic lines (e.g., Charbonneau et al. 2002; Vidal-Madjar et al. 2003; Redfield et al. 2008; Jensen et al. 2012; Hoeijmakers et al. 2018; Nortmann et al. 2018; Casasayas-Barris et al. 2019; Cauley et al. 2019; Deibert et al. 2019) and molecular species (e.g., Snellen et al. 2010; Brogi et al. 2012; Kreidberg et al. 2015; Sing et al. 2016; Turner et al. 2016b, 2017; Birkby et al. 2017; Esteves et al. 2017; Evans et al. 2018) in the planetary atmosphere. Additionally, extended hydrogen envelopes have been observed around several hot Jupiters (e.g., Vidal-Madjar et al. 2003; Lecavelier des Etangs et al. 2012). Measuring the slope of the transmission spectra can be used to infer the presence of atmospheric clouds or haze (e.g., Pont et al. 2008; Bean et al. 2010; Turner et al. 2013, 2016b; Knutson et al. 2014; Kreidberg et al. 2014; Sing et al. 2016) and scattering particles (e.g., Pont et al. 2008; Sing et al. 2011; Turner et al. 2016b; Mallonn & Wakeford 2017).

KELT-9b orbits an A0V/B9V star at a distance of 0.034 au (Gaudi et al. 2017) and is the hottest known gas giant with a

 $T_{\rm eff} \approx 4500$ K (Hooton et al. 2018; Wong et al. 2019; Mansfield et al. 2019). Due to the high temperature of the planet Kitzmann et al. (2018) predicted that iron absorption lines and several near-ultraviolet resonance lines should be observable in KELT-9b's transmission spectra. Neutral and ionized iron in its atmosphere have subsequently been detected in multiple HARPS-N observations (Hoeijmakers et al. 2018, 2019). Also, Yan & Henning (2018) reported an extended hydrogen atmosphere (~1.64 R_p) using H α observations with the CARMENES spectrograph. The H α detection was confirmed by Cauley et al. (2019), who observed variability in the in-transit absorption light curve suggesting that the transiting gas is not uniform. Using high-resolution optical spectra from HARPS-N during four separate KELT-9b transits Borsa et al. (2019) observed the atmospheric Rossiter-McLaughlin effect. In total, 14 atomic species have been reported previously in KELT-9b's atmosphere (Hoeijmakers et al. 2018, 2019; Yan & Henning 2018; Cauley et al. 2019).

Here, we present observations of KELT-9b using the CARMENES spectrograph on the 3.5 m Calar Alto telescope. With these data, we detect ionized calcium (Ca II) for the first time in its atmosphere. Ca II is predicted to be in hot Jupiter atmospheres (e.g., Turner et al. 2016a) and has been observed in the atmospheres of the super-Earth 55 Cnc e (Ridden-Harper et al. 2016) and the hot Jupiter KELT-20b (Casasayas-Barris et al. 2019). Our data also reveal prominent H α absorption, confirming the presence of an extended hydrogen envelope (Yan & Henning 2018 and Cauley et al. 2019).

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Figure 1. Example of the analysis steps done on the CARMENES data. (a) Reduced data from the CARACAL v2.01 pipeline. (b) Data after blaze correction. (c) Data after blaze correction and the telluric and stellar lines and other systematics have been removed using SYSREM.

2. Observations and Data Reduction

We observed the transit of KELT-9b on 2018 June 6–7 from 21:34–3:53 UT with the CARMENES instrument (Quirrenbach et al. 2016) on the 3.5 m Calar Alto telescope. For the main analysis in this Letter, we use data from the visible spectrograph that covers the wavelength range 5200–9700 Å, over 61 spectral orders with a resolution of ~94,600. To search for helium we also use the near-infrared (near-IR) spectrograph that has a wavelength range of 9600–17100 Å, over 28 spectral orders at a resolution of ~80,400. Exposure times of 113 and 121 s were used for the optical and near-IR spectrographs, respectively. Over the entire observation 140 exposures were taken. The airmass varied from 2.10 to 1.07 during the night.

Standard data reduction was performed with the CAR-MENES Reduction and Calibration (CARACAL v2.01) pipeline (Caballero et al. 2016). This includes bias correction, flat-fielding, cosmic-ray removal, and wavelength calibration. An example of the result of this standard reduction can be found Figure in 1(a). An additional step was taken to ensure more accurate wavelength calibration. We used a telluric transmission spectrum taken from the ESO Cerro Paranal Advanced Sky Model (Moehler et al. 2014) and crosscorrelated it with each order of our data. The wavelengths of each order were then linearly shifted by the value that optimized the cross-correlation. We found offsets between 0.13 and 0.2 Å depending on the order. Next, we performed a blaze correction (Section 2.1) on the data and then removed the telluric and stellar lines and other systematics from the data using SYSREM (Section 2.2).

2.1. Blaze Normalization

As a first step, we correct each order individually by its blaze function. To start, we divide the first science spectrum by the rest of the frames in order to divide out time-invariant changes. Next, each order in this normalized data was binned where outliers in each bin were masked out using the PATROL sigma thresholding technique (Turner et al. 2019) and these binned data were fit to a polynomial. Finally, we divided each order in the data from the CARACAL v2.01 pipeline by this polynomial fit to remove the blaze function. An example of the blaze-corrected data can be found in Figure 1(b).

2.2. Telluric and Stellar Subtraction Using SYSREM

Next, the telluric and stellar lines and other systematics were removed using SYSREM (Tamuz et al. 2005) from the blazecorrected data. This algorithm has been used extensively to study exoplanet atmospheres (e.g., Birkby et al. 2017; Esteves et al. 2017; Deibert et al. 2019). SYSREM interactively identifies and subtracts systematics (e.g., telluric lines) that are time-invariant while preserving the Doppler-shifted exoplanet signal. Five iterations of SYSREM were ran where each order was treated as a separate "light curve." If we run between one and five iterations with SYSREM, we obtain the same results within 1σ for the absorption profiles and subsequent physical information (see Figure 4(a) in Appendix A). Birkby et al. (2017) showed that a small part of the planet signal is removed after a few iterations of SYSREM. Therefore, to minimize this effect we use the number of iterations that maximized the signal-to-noise ratio in our detection. An example of the data after SYSREM was performed can be found in Figure 1(c).

To check the reliability of the data reduction with SYSREM, we also corrected the data using the telluric reduction method described in Brogi et al. (2012) and in Appendix A. Using this method, we find consistent results within 1σ with SYSREM for the absorption line profiles and widths. A comparison of the H α and Ca II absorption profiles produced using both methods can be found in Figures 4(b)–(c) in Appendix A.

2.3. Correcting for the Rossiter-McLaughlin Effect

Finally, we removed the stellar Rossiter–McLaughlin (RM) effect. We modeled these effects using the technique outlined in de Mooij et al. (2017) and detailed below. This model was subtracted from the data to get the final spectrum for each order.

The stellar disk was modeled using a regular grid and has a radius of 510 pixels. The stellar spectrum used for each pixel was obtained from the Spectroscopy Made Easy (SME) software package (Piskunov & Valenti 2017). The stellar parameters used for KELT-9 in the model are $T_{\rm eff} = 10,170$ K, $\log(g) = 4.093$, [Fe/H] = -0.03, and $V \sin I = 111.4$ km s⁻ (Gaudi et al. 2017). The spectra were generated at 25 different limb angles ($\mu = \cos \theta$) and linearly interpolated in μ for all other limb angles, to ensure that our model takes the (differential) limb-darkening effects into account. We assume that the star rotates as a solid body, and ignore the effects of gravity darkening. The planet was modeled as an opaque disk in a circular orbit with the planetary parameters (radius, semimajor axis, orbital period, impact parameter, and projected spin-orbit misalignment angle) taken from Gaudi et al. (2017). We note that when modeling the RM effect, we use the R_p/R_* from Gaudi et al. (2017), and do not consider the impact of the excess absorption in the planet's atmosphere at different wavelengths. This will lead to a slight underestimate of the RM amplitude in lines with strong planetary absorption, and thereby a small underestimate of the planetary absorption. After generating the models for each of the frames, we normalized them, and subsequently removed the out-of-transit model in order to obtain the RM effect. We note that it is crucial to normalize the models, as the data are effectively continuum normalized by the blaze correction. This normalization will lead to a slight increase in the depth of the line for the parts not occulted by the planet.

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3. Results and Discussion

We searched for absorption due to individual spectral lines using the final dynamic spectrum after blaze normalization, removing the telluric and stellar lines, and correcting for the RM effect. The composite absorption spectra were generated by correcting for the systemic velocity of the system and then shifting each individual observation into the rest frame of the planet. We used the planetary parameters from Gaudi et al. (2017), setting the radial velocity semi-amplitude $K_p = 276$ m s⁻¹ and the systemic velocity $V_r = -19.819$. We note that the K_p values in the literature differ by as much as 5σ , so care must be taken when looking for offsets in line profiles due to atmospheric winds. To obtain the final composite absorption profiles, we averaged all the data between the second and third transit contact points excluding the region affected by the Doppler shadow.

Using these spectra, we found absorption due to H α and the ionized calcium (Ca II) infrared triplet Doppler-shifted with the planet's orbital motion (Figure 2; Table 1). For Ca II, the planetary absorption is clearly seen as Doppler-shifted absorption in the spectra at 8500.35 Å line and marginally in the 8544.44 Å line. The planetary absorption profiles for the H α and Ca II can be found in Figure 2. The line profiles for the Ca II lines have been rebinned by two data points to enhance the significance of the detection. H α and all three lines of the Ca II infrared triplet are seen clearly in the absorption profiles.

The planetary line transmission spectra (T_{ν}) are modeled as a Gaussian function

$$T(\nu) = 1 + \delta e^{\frac{(\nu - \nu_{cen})^2}{2\sigma^2}},$$
 (1)

where δ is the planetary absorption depth (formally the Gaussian height), ν is the velocity in the planetary rest frame, ν_{cen} is the center velocity of the absorption profile, and σ is the Gaussian standard deviation (FWHM = 2.355 σ). The best-fit model parameters (h, ν_{cent} , and σ) and their errors were found by performing a Differential Evolution Markov Chain Monte Carlo (DE-MCMC) analysis. The same DE-MCMC code is used in the transit modeling program EXOMOP (Turner et al. 2016b). In all, we ran 20 chains and 10⁶ links and used the Gelman–Rubin statistic to guarantee chain convergence. The results of the modeling can be found in Table 1.

Using this modeling, we find a line absorption depth for H α of 1.02 \pm 0.09%, where the optical transit depth of KELT-9b is 0.68% (Gaudi et al. 2017). The hydrogen atmosphere corresponds to a radius of ~1.58 R_p , which extends near the Roche-lobe radius of 1.91 R_p (Yan & Henning 2018). Our H α line depth is within 1 σ with those found by Yan & Henning (2018) and Cauley et al. (2019) of 1.15 \pm 0.05% and 1.103 \pm 0.010%, respectively. The consistency suggests that the hydrogen envelope around KELT-9b is relativity stable over several years.

Next, we will estimate the hydrogen column density that caused the observed transit signal. Using the analytic equations from Huang et al. (2017), the number density in the hydrogen 2ℓ state $(n_{2\ell})$ can be estimated using the line center optical depth (τ_0)

$$\tau_0 = 35 \left(\frac{n_{2\ell}}{10^4 \text{cm}^{-3}} \right). \tag{2}$$

Furthermore, τ_0 can be found by measuring the spectral line width $(\Delta \nu)$ of H α

$$\Delta \nu = 9.1 \,\mathrm{km} \,\mathrm{s}^{-1} (\ln \tau_0)^{1/2} \,. \tag{3}$$

Therefore, directly measuring $\Delta\nu$ from the data can provide an estimate on the number density. We fit a line width (σ in the MCMC fit) of $16.3^{+1.7}_{-1.5}$ km s⁻¹. The line width is mainly affected by thermal broadening and the optical depth at the line center. A line width larger than the thermal width indicates that the gas is optically thick. The effect of thermal broadening on the line width ($\Delta\nu_T$) can be measured as

$$\Delta \nu_T = \sqrt{\frac{8k \ln 2}{mc^2}} T^{1/2} f_0, \qquad (4)$$

where *T* is the temperature and f_0 is the oscillator strength. We find a thermal width of 9.1 km s⁻¹ for the H α line where $f_0 = 0.64108$ and assuming a temperature of 5000 K. Huang et al. (2017) calculated a mean temperature of 5000 K in the atomic layer of their atmospheric model by balancing photoelectric heating and atomic line cooling. An optical depth of $\tau_0 \approx 25$ and a number of density of $n_{2\ell} = 7.1 \times 10^3 \text{cm}^{-3}$ are found using the measured line width. Therefore, our data confirm that the the H α line is optically thick.

We find absorption depths of 0.40 ± 0.05 , 0.51 ± 0.05 , and $0.48 \pm 0.06\%$ for the Ca II 8500.35 Å, 8544.44 Å, and 8664.52 Å, lines, respectively (Figure 2; Table 1). Using Equation (3) and the measured line widths, we find $\tau_0 \approx 1.3$. Thus, the Ca II lines are optically thick. Our observations confirm the predictions by Turner et al. (2016a) that Ca II should be observable in the upper atmospheres of hot Jupiters. Additionally, our Ca II detection was confirmed recently using independent CARMENES and HARPS-N observations by Yan et al. (2019). The implications of the Ca II detection are discussed further in Section 3.1.

Helium (He I; 10830 Å) has been predicted to exist in hot Jupiter atmospheres (Turner et al. 2016a; Oklopčić & Hirata 2018) and has been detected subsequently (e.g., Allart et al. 2018, 2019; Nortmann et al. 2018; Spake et al. 2018). We searched our data for He I absorption and found none, in agreement with the results of Nortmann et al. (2018). Theoretical modeling suggests that, for planets orbiting A stars, photoionization can easily deplete the fraction of helium atoms in the triplet state (Oklopčić 2019), thus our nondetection may not be surprising.

3.1. Model of KELT-9b's Atmosphere

In addition to the parametric fits, we have modeled the observed line profiles with synthetic NLTE transmission spectra produced using the spectral synthesis code Cloudy (Ferland et al. 2017). For the computation of the transmission spectra, we assume a spherically symmetric planetary atmosphere that varies only with altitude. The atmospheric abundance ratio profiles of the first 30 elements, including ions and molecules, as well as the temperature and pressure profiles are adopted from a 1D PHOENIX model atmosphere as presented in Lothringer et al. (2018) and Fossati et al. (2018). All other planetary and orbital properties are taken from Gaudi et al. (2017). Specifically, the mass of the planet is found to be $2.88 \pm 0.84 M_{Jup}$. We compute the radius at each layer by first determining the pressure scale



Figure 2. Spectra of the ultra-hot Jupiter KELT-9b from the CARMENES spectrograph on the 3.5 m Calar Alto telescope during transit are shown surrounding the H α and ionized calcium (Ca II) infrared triplet lines. The left column shows the final spectrum with absorption from the planet's atmosphere, Doppler-shifted due to its changing radial velocity during transit (green dashed line). The ingress and egress of the transit are shown as solid black lines. The right column shows the H α and Ca II composite absorption profiles in the rest frame of KELT-9b. The red dotted–dashed line is a Gaussian fit to the data and the blue dashed curve and green dashed curve are a synthetic atmospheric model with a planetary mass of 2.88 M_{Jup} and 2.08 M_{Jup} , respectively (see Section 3.1 for more details). The depth of the H α line is 1.02 ± 0.08% (compared to the optical transit depth of 0.67%) and indicates the existence of an extended hydrogen envelope around the planet. We find absorption depths of 0.40 ± 0.5, 0.51 ± 0.51 ± 0.5, and 0.48 ± 0.6% for the Ca II 8500.35 Å, 8544.44 Å, and 8664.52 Å lines, respectively.

Turner et al.

 Table 1

 Summary of Atmosphere Line Absorption Parameters

Species	Wavelength (Å)	Depth (%) (%)	$(\mathrm{km \ s}^{-1})$	σ (km s ⁻¹)	FWHM (km s ⁻¹)
Ηα	6564.59 (6562.79)	1.02 ± 0.09	$-2.10^{+1.79}_{-1.82}$	$16.3^{+1.7}_{-1.5}$	$38.3^{+4.0}_{-3.5}$
Ca II	8500.35 (8498.02)	0.40 ± 0.05	$-1.76^{+1.47}_{-1.43}$	$9.8^{+1.3}_{-1.1}$	$23.1^{+3.0}_{-2.7}$
Ca II	8544.44 (8542.09)	0.51 ± 0.05	$1.89^{+1.24}_{-1.38}$	$10.0^{+1.1}_{-1.3}$	$23.6^{+2.7}_{-3.1}$
Ca II	8664.52 (8662.14)	0.48 ± 0.06	$1.99^{+1.34}_{-1.44}$	$8.3^{+1.1}_{-1.0}$	$19.6^{+2.6}_{-2.4}$

Note. Column 1: species detected. Column 2: vacuum wavelength of the species (air wavelength). Column 3: transit depth of the absorption profile. Column 4: center velocity of the absorption profile in the rest frame of the planet. Column 5: standard deviation derived from the the absorption profile assuming a Gaussian profile. Column 6: FWHM of the absorption profile derived using σ assuming a Gaussian profile.

Table 2							
Summary of Model Atmospheric Parameters at Line Formation Altit	tudes						

Species	Effective Radius (R_p)	Pressure (nbar)	$n \pmod{(\mathrm{cm}^{-3})}$	σ (cm ⁻²)
Ηα	1.28	7.9	$3.6 \times 10^{1}(2s)$	$1.8 \times 10^{12}(2s)$
			$1.1 \times 10^{2}(2p)$	$5.2 \times 10^{12}(2p)$
$H\alpha$ (low mass)	1.44	6.3	$2.8 \times 10^{1}(2s)$	$2.3 \times 10^{12}(2s)$
			$8.6 \times 10^{1}(2p)$	$6.9 \times 10^{12}(2p)$
Ca II 8500	1.21	61	9.8×10^{3}	4.5×10^{13}
Ca II 8500 (low mass)	1.32	50	6.7×10^{3}	4.2×10^{13}
Ca II 8544	1.26	14	6.5×10^{2}	3.3×10^{12}
Ca II 8544 (low mass)	1.40	13	1.4×10^{2}	4.1×10^{12}
Ca II 8664	1.24	20	1.4×10^{3}	6.9×10^{12}
Ca II 8664 (low mass)	1.38	17	1.0×10^3	6.7×10^{12}

Note. Column 1: species. Column 2: effective planetary radius at line center. Column 3: atmospheric pressure at effective radius. Column 4: species number density at effective radius. Column 5: species line-of-sight column density at effective radius. For H α densities, values listed are for n = 2 excitation stages.

height,

$$H_i = \frac{R T_i}{\mu_i g_i},\tag{5}$$

where *R* is the gas constant, T_i is the temperature of the *i*th layer of the atmosphere, μ_i is the mean molecular weight of the *i*th layer, and g_i is the planetary gravity evaluated at the *i*th layer (assuming constant planetary mass), and inserting it into

$$r_{i} = \begin{cases} -H_{i} \ln(P_{i}/P_{i-1}) + r_{i-1} & P_{i} > P_{0} \\ -H_{i} \ln(P_{i}/P_{i+1}) + r_{i+1} & P_{i} < P_{0} \end{cases}$$
(6)

where H_i is the pressure scale height evaluated at the *i*th layer, P_i is the atmospheric pressure of the the *i*th layer, and P_0 is the reference pressure at the planetary surface radius, $r_0 = R_P$. We take P_0 to be $P_0 = 0.01$ bar, the pressure in the PHOENIX model at which the equilibrium temperature occurs. This formulation neglects to account for Roche potential effects, which may be important for KELT-9b and will need to be taken into account for more detailed calculations extracting physical information on the properties of the planetary atmosphere from the data.

We map the 1D model properties onto concentric circles, and calculate the lengths through successive layers of atmosphere along line-of-sight transmission chords as

$$l_i = c_i - c_{i-1} = \sqrt{2r_i h_i - h_i^2} - \sqrt{2r_{i-1} h_{i-1} - h_{i-1}^2}, \quad (7)$$

where l_i is the length through layer *i*, c_i is half the chord length along line of sight through layer *i*, r_i is the radius of layer *i*, and

 h_i is the height down from the top of the atmosphere to layer *i*. These lengths, along with the atmospheric properties of their respective layers, are stacked and entered into Cloudy as the line-of-sight transmission medium. A schematic of this setup is shown in Figure 5 in Appendix B.

We compute separate transmission spectra with Cloudy for each layer, at a spectral resolution of R = 120,000. The total transmission spectrum of the planet is then the sum of these layer spectra, weighted by the relative area of the stellar disk covered by the respective layers along each line of sight. We assume a nonlinear limb-darkening function in which we employed limb-darkening coefficients extracted on the basis of a library of PHOENIX models (Husser et al. 2013) as in Salz et al. (2019), and average the wavelength dependence over a 10 Å window centered on each line of interest. A more detailed description of the computation of the synthetic transmission spectrum is forthcoming in M. E. Young et al. (2019, in preparation).

The right column in Figure 2 displays the observed data in comparison to our synthetic profiles. For H α and the Ca II triplet lines, we find that the observed line width is well fit by the synthetic spectra accounting for a turbulent velocity of 3 km s⁻¹, but the line depth is significantly underestimated. The mismatch may be due to the temperature in the model being too low at the line formation layers. We roughly emulate the effect of a higher temperature by reducing the planetary mass in our spectral synthesis by $0.8 M_{Jup}$ (2.08 M_{Jup} , a mass that is still within 1 σ of the derived mass of the planet), thus inflating the atmosphere via reducing the gravity in the scale height calculation. This lower-mass transmission spectrum has an increased line depth and fits the data.



Figure 3. Atmospheric profiles for the H α and the Ca II infrared triplet lines derived from the atmospheric model. (a) Pressure-column density profile of H α . H α is formed at a pressure of ~6 nbar with a number density of 28 cm⁻³ for the 2s state and 86 cm⁻³ for the 2p state. (b) Pressure-density profile of three Ca II infrared lines. The lines are formed between 13 and 50 nbar with number densities between 0.6 and 6.7 \times 10³ cm⁻³.

Table 2 lists the atmospheric altitude, pressure, and temperature obtained from the PHOENIX model at which the peak of the H α and Ca II triplet lines form considering both models (i.e., planetary mass of 2.08 and 2.88 M_{Jup}). The H α line forms at a pressure of about 6 nbar while the Ca II lines form slightly deeper, at pressures ranging between 13 and 50 nbar. Figure 3 shows the atmospheric density profiles with the pressures at which each line forms.

The 30% decrease in planetary mass corresponds to a \sim 1900 K increase in atmospheric temperature, implying that the temperature between the 1 and 100 nbar level should be about 8000 K. Therefore, the observations indicate that the atmosphere should be hotter than predicted in the PHOENIX model. However, this does not include the effect that a hotter atmosphere would have on the level populations involved in the line formation, which would probably lead to a further increase in the line strength, hence requiring a slight decrease in the atmospheric temperature. We can therefore conclude that the atmospheric temperature between 10 and 100 nbar should lie between 6100 and 8000 K. Our finding is in agreement with García Muñoz & Schneider (2019), who also found the need for a lower planetary mass, hence higher atmospheric temperature, to fit the observed H α line profile.

The peak of the H α line, the one forming at the lowest pressure (hence higher altitude) among the four, forms at an altitude of about 1.44 planetary radii (see Table 2), which is

about 0.5 planetary radii below the planetary Roche lobe $(\sim 1.91 R_p)$ at the terminator. This shows that the atmosphere at the levels probed by the H α line is not in a hydrodynamic regime and that on KELT-9b this line does not directly probe the escaping upper atmosphere. Instead, the H α line, as well as the other observed Balmer and metal lines, act as atmospheric thermometers enabling one to probe the temperature profile, thus the energy budget, which is a key element in controlling the escape. In fact, it is still believed that KELT-9b hosts an escaping upper atmosphere (Fossati et al. 2018; García Muñoz & Schneider 2019), which therefore will require ultraviolet observations to be directly probed (e.g., Fossati et al. 2010; Sing et al. 2019).

4. Conclusions

In this study, we present high-resolution $(R \sim 94,600)$ transit observations of KELT-9b using the CARMENES spectrograph on the 3.5 m Calar Alto telescope. We unambiguously detect atmospheric absorption due to $H\alpha$ and the Ca II infrared triplet Doppler-shifted with the planet's orbital motion (Figure 2; Table 1). These detections are very robust because different methods used to remove telluric and stellar lines from the data obtained the same final results (Section 2.2; Appendix A). This detection of Ca II is the first time this species has been observed in KELT-9b's atmosphere and adds to the suite of the 14 already known atmospheric constituents (Hoeijmakers et al. 2018, 2019; Yan & Henning 2018; Cauley et al. 2019). All four absorption lines are found to be optically thick. We also fit the observed absorption line profiles with synthetic NLTE transmission spectra produced using the spectral synthesis code Cloudy (Section 3.1; Appendix B). From this model, we derived the atmospheric altitude, pressure, temperature, and number density where the H α and Ca II triplet lines are formed (Table 2; Figure 3). The Ca II lines are formed at pressures between 13 and 50 nbar corresponding to an effective radius between 1.32 and 1.40 R_p . The H α line forms at an effective radius of 1.44 R_p corresponding to a pressure of \sim 6 nbar. We also find that the atmospheric temperature between 10 and 100 nbar should lie between 6100 and 8000 K. Since the H α line is formed well within the planetary Roche lobe (1.91 R_p ; Yan et al. 2017), our modeling demonstrates for the first time that on KELT-9b the H α line does not probe the escaping upper atmosphere, though it constrains the atmospheric energy budget, which is an important element controlling the escape. In this Letter, we show the promise of using individual absorption lines to study the structure of exoplanetary atmospheres.

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Figure 4. (a) Here we show the H α absorption profiles in the rest frame of KELT-9b derived using a different number of iterations in SYSREM. The error bars are derived from the observational data and only one set is shown for clarity. All of the two profiles are identical within the 1 σ error bars. (b) and (c) Here we show the H α and Ca II absorption profiles in the rest frame of KELT-9b derived using SYSREM and the method described in Brogi et al. (2012). The error bars are derived from the observational data cube, and we only show one set of error bars for clarity. The two profiles are nearly identical within the 1 σ error bars.

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Facilities: 3.5 m Calar Alto telescope/CARMENES, Exoplanet Archive.

Software: CARMENES Reduction and Calibration (CAR-ACAL v2.01) pipeline (Caballero et al. 2016); IDL Astronomy Users Library; Cloudy (Ferland et al. 2017); Spectroscopy Made Easy (SME) software package (Piskunov & Valenti 2017); we used the following Python packages: Astropy, SciPy, NumPy.

Appendix A SYSREM Tests

In order to efficiently run SYSREM, the procedure has to be performed multiple times. Each iteration removes a different order of systemics due to the telluric and stellar lines. To confirm that running multiple iterations of SYSREM is not removing part of the exoplanet signal we overplot the absorption profiles for H α for different SYSREM iterations in Figure 4(a). All the profiles are identical within the 1σ error bars.

To check the reliability our data reduction with SYSREM, we compared our results to those obtained by using the telluric reduction method described in Brogi et al. (2012) and briefly below. First, the flux of each wavelength is fit as a linear function of the geometric airmass at the time of each observation. Next, the flux of a few of the deepest telluric lines were measured over time and the rest of the wavelengths were corrected using a linear regression. Finally, a high-pass filter is applied to each wavelength to remove any remaining low-order structure. Figure 4 shows the comparison of the H α (panel b) and Ca II (panel c) absorption profiles between the two methods. We find consistent results within 1σ for both methods. Therefore, if SYSREM is removing part of the planet signal in our data the amount removed is inconsequential on our observed profiles and the resulting physical information of the planet.

Appendix B Schematic of Cloudy Model

A schematic of the atmospheric setup in the cloudy model presented in Section 3 can be found in Figure 5.



Figure 5. Schematic diagram of transmission chords along the line of sight for transmission spectrum calculation. c is half the chord length along line of sight, l is the path length through a given layer of atmosphere, h is the height down from the top of the atmosphere to the layer, r is the radius of the layer, and the subscripts denote which layer.

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